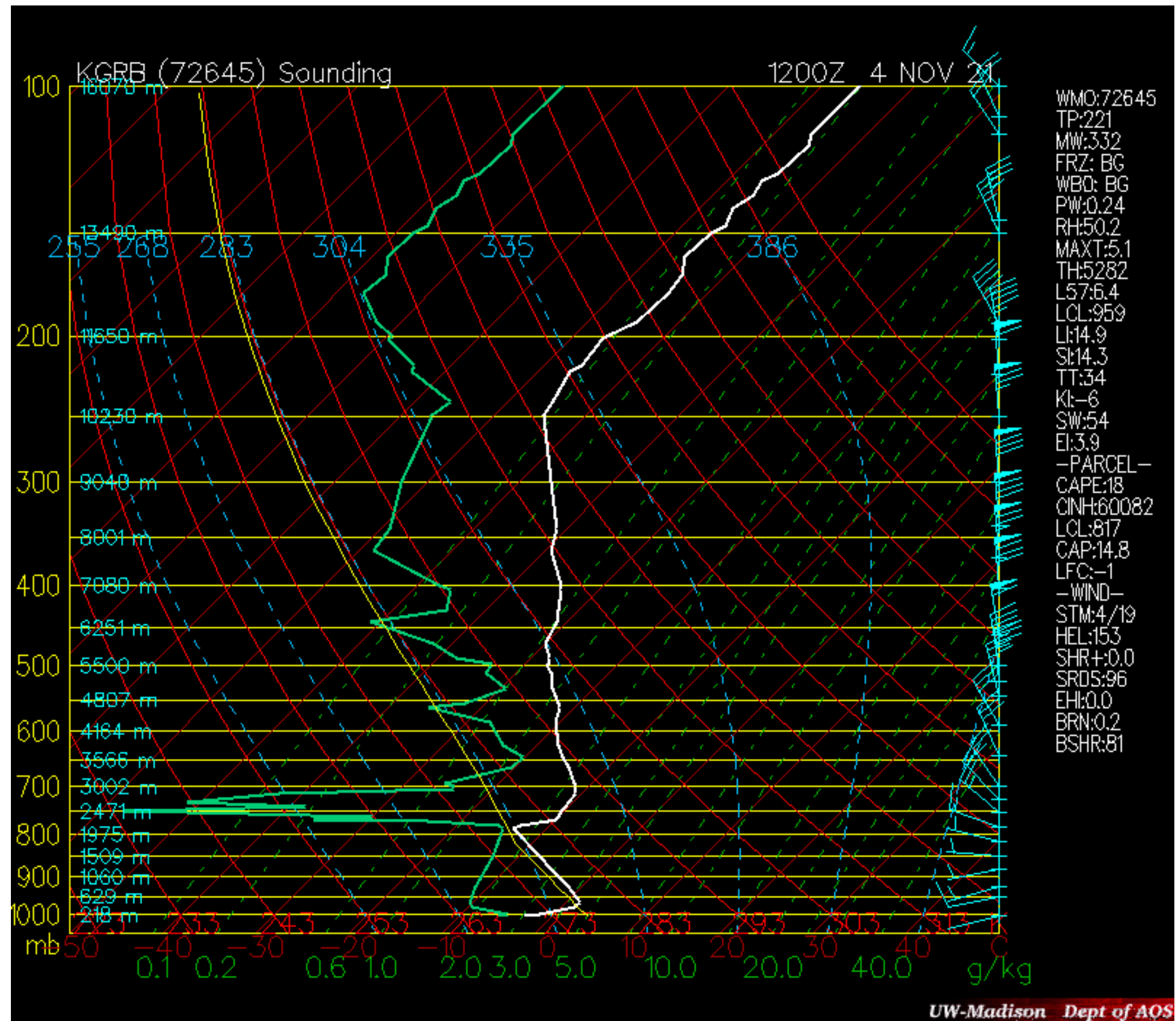


AOS 630: Introduction to Atmospheric
and Oceanic Physics
Lecture 17 Fall 2021
*Carnot Engine 2,
Buoyancy and convection*

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Next Tuesday we will discuss the section titled “**The mature hurricane: A natural Carnot engine**” by Emanuel (1991) (TC_Carnot_Engine.pdf on Canvas).



[https://earth.nullschool.net/#2021/08/25/1600Z/
 wind/isobaric/500hPa/
 orthographic=-87.74,30.76,1229](https://earth.nullschool.net/#2021/08/25/1600Z/wind/isobaric/500hPa/orthographic=-87.74,30.76,1229)

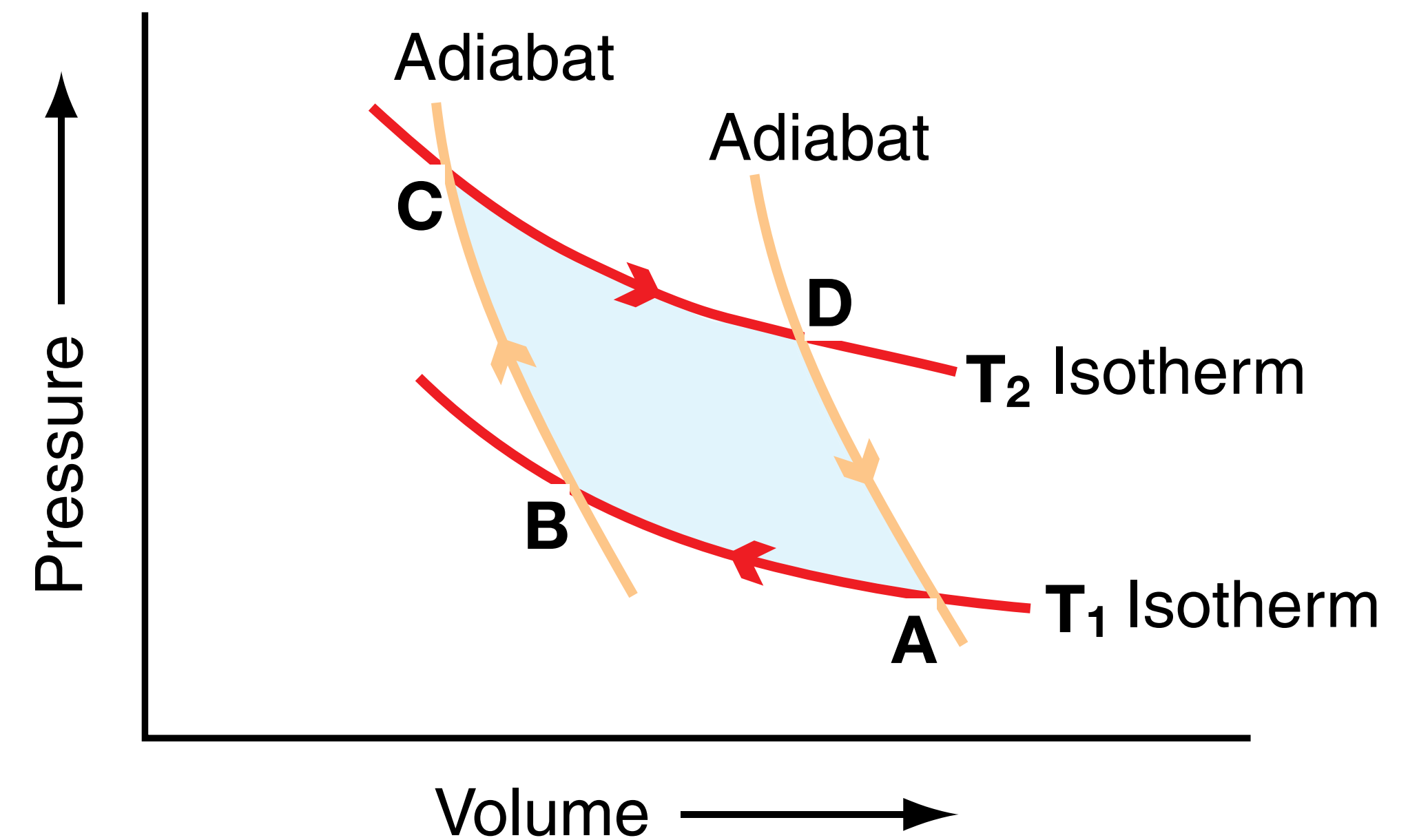
Last Class: Carnot Engine

The first law (internal energy form) integrated over this cycle takes the form:

$$\oint c_v dT = \oint \delta q - \oint \delta w$$

Because state variables don't change during a closed loop integral, it follows that

$$\oint \delta q = \oint \delta w$$



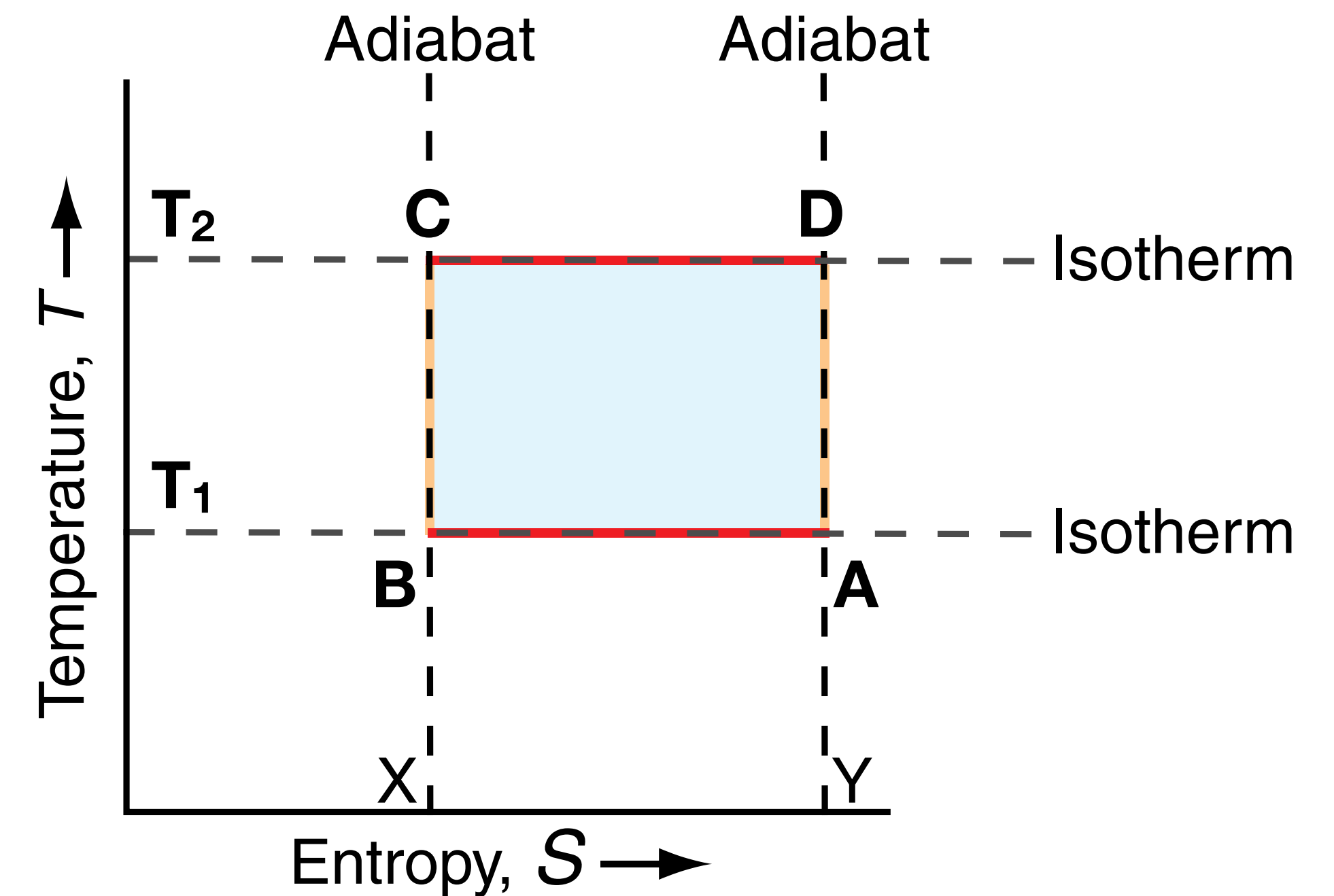
Carnot Engine

Writing these in exact derivative form we have

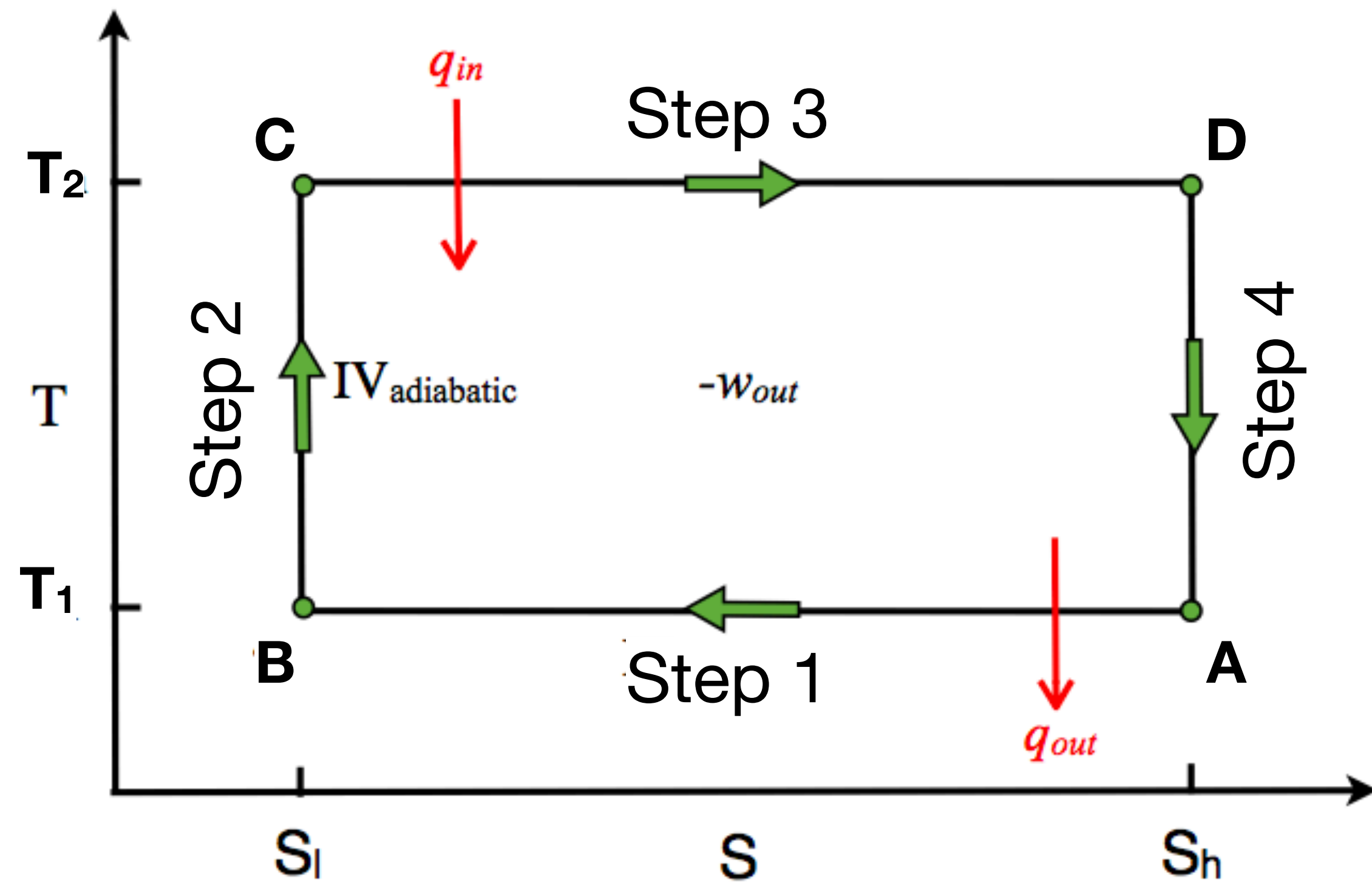
$$\oint T ds = \oint p d\alpha \neq 0$$

Now let's consider a cycle divided into 4 steps:

1. Isothermal compression at a cooler T_1
2. Adiabatic compression to T_2
3. Isothermal expansion at T_2
4. Adiabatic expansion back to T_1

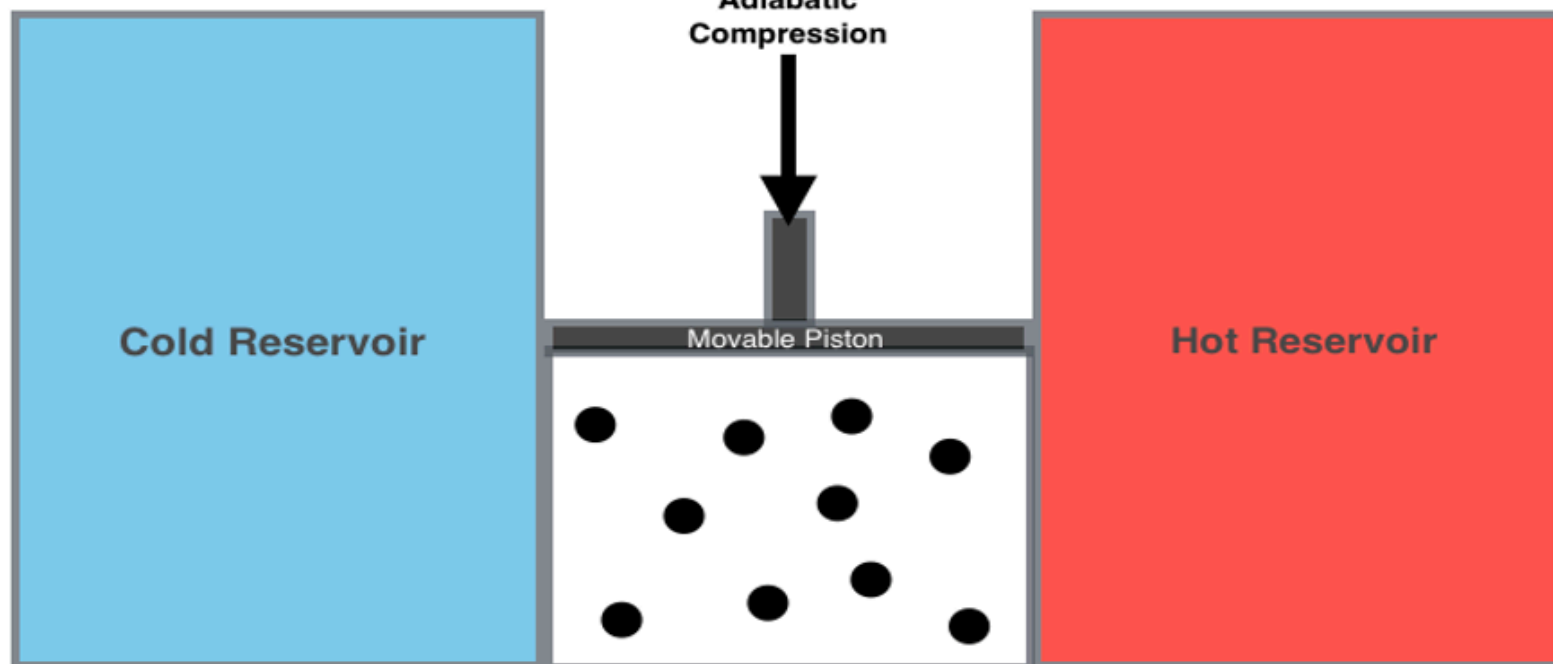
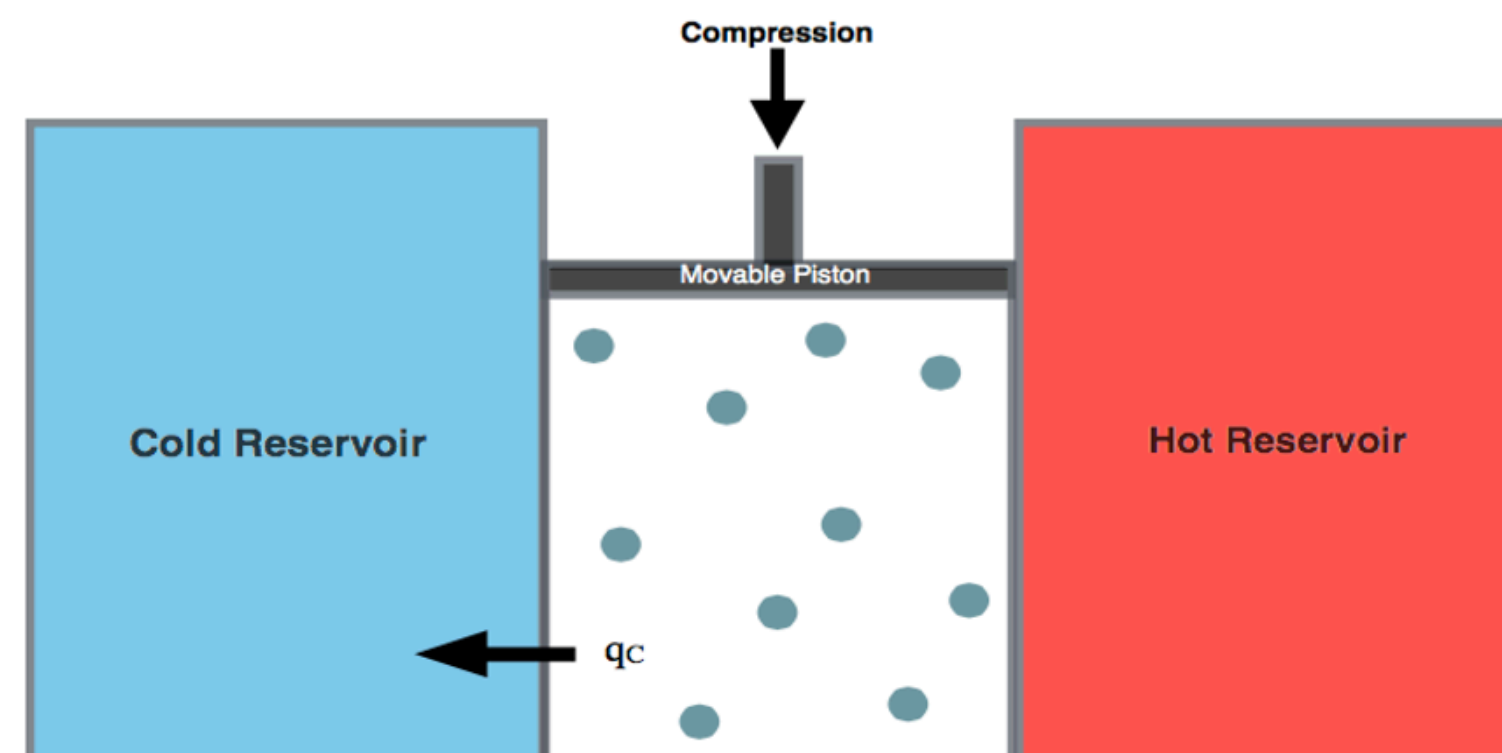
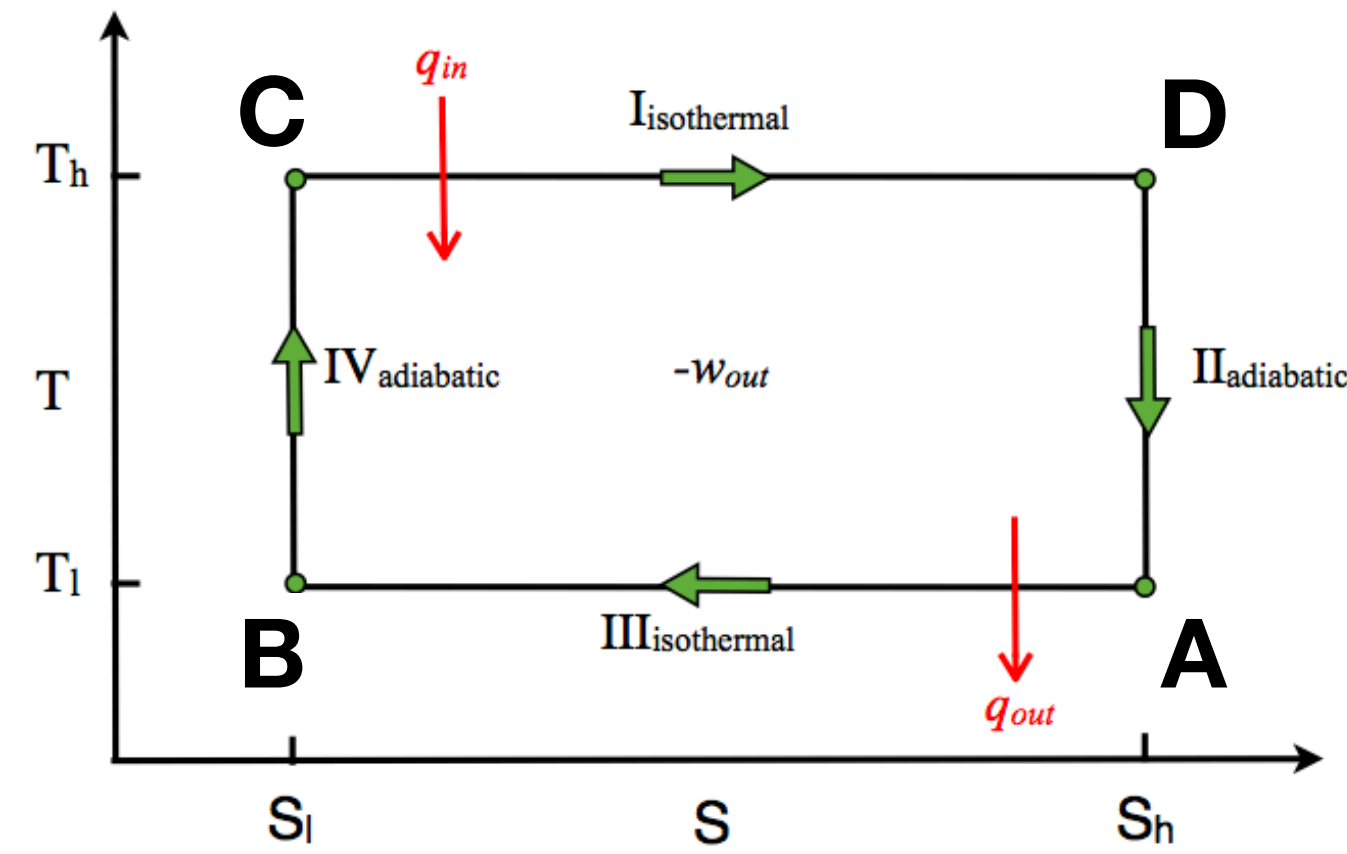
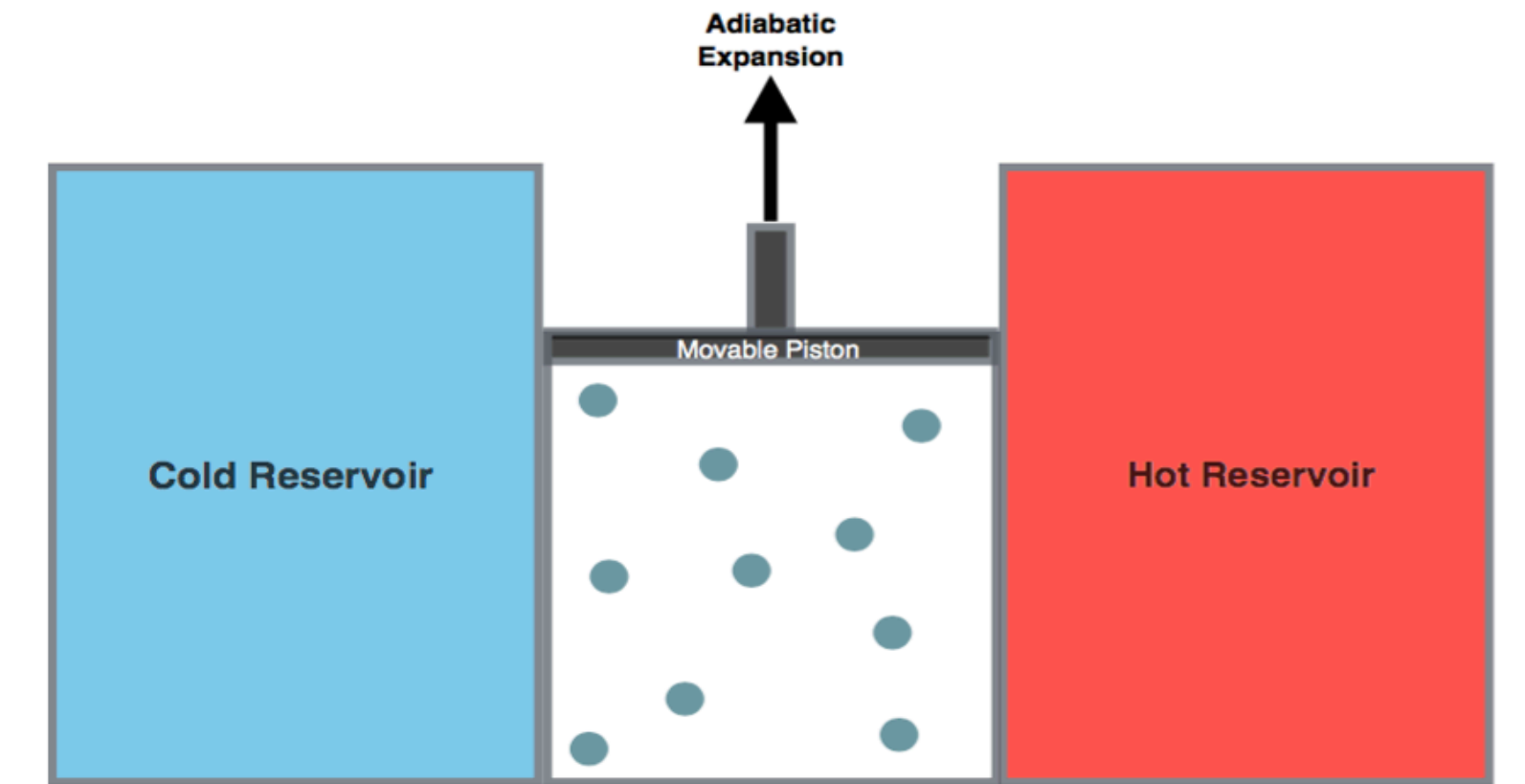
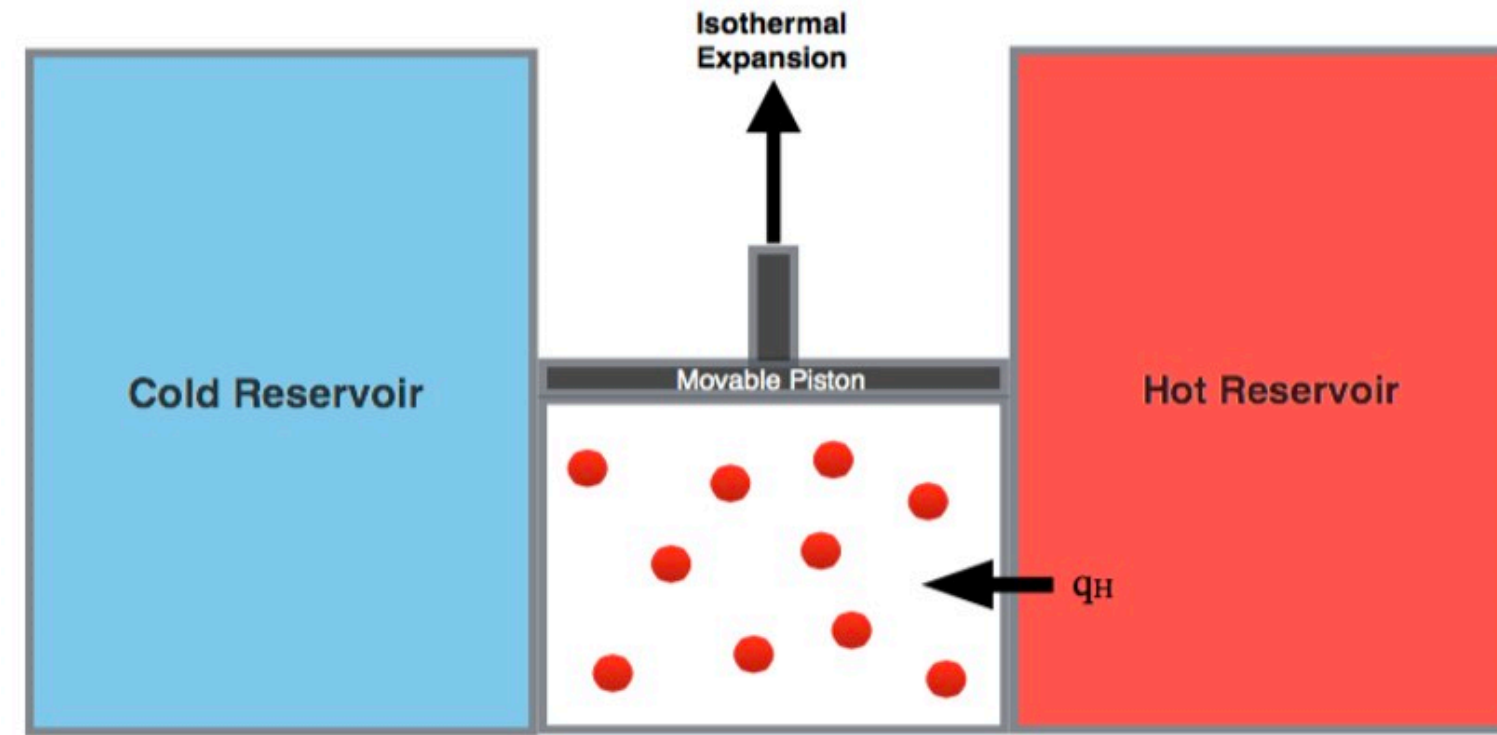


Carnot Engine



1. Isothermal compression at a cooler T_1
2. Adiabatic compression to T_2
3. Isothermal expansion at T_2
4. Adiabatic expansion back to T_1

Carnot Engine



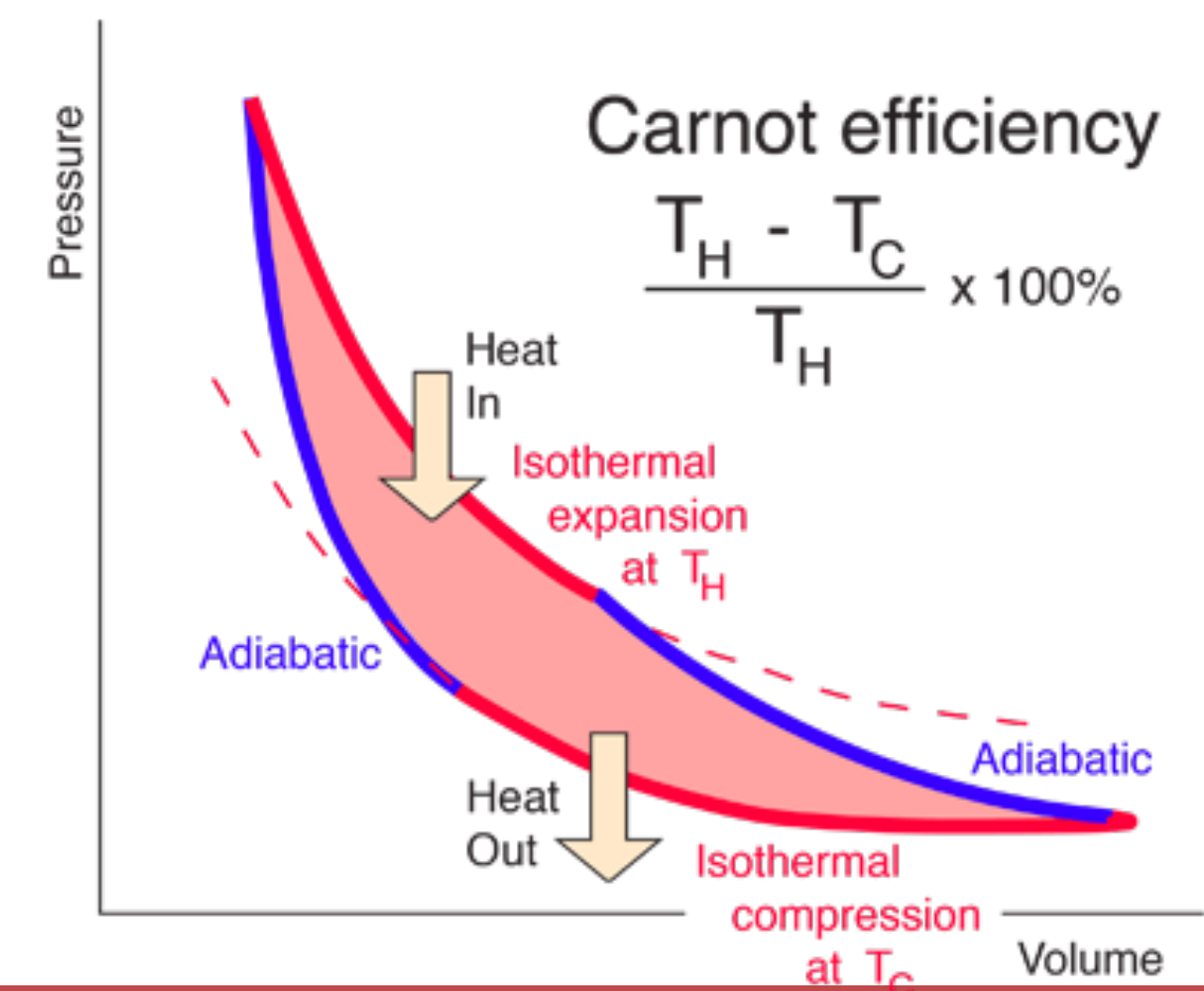
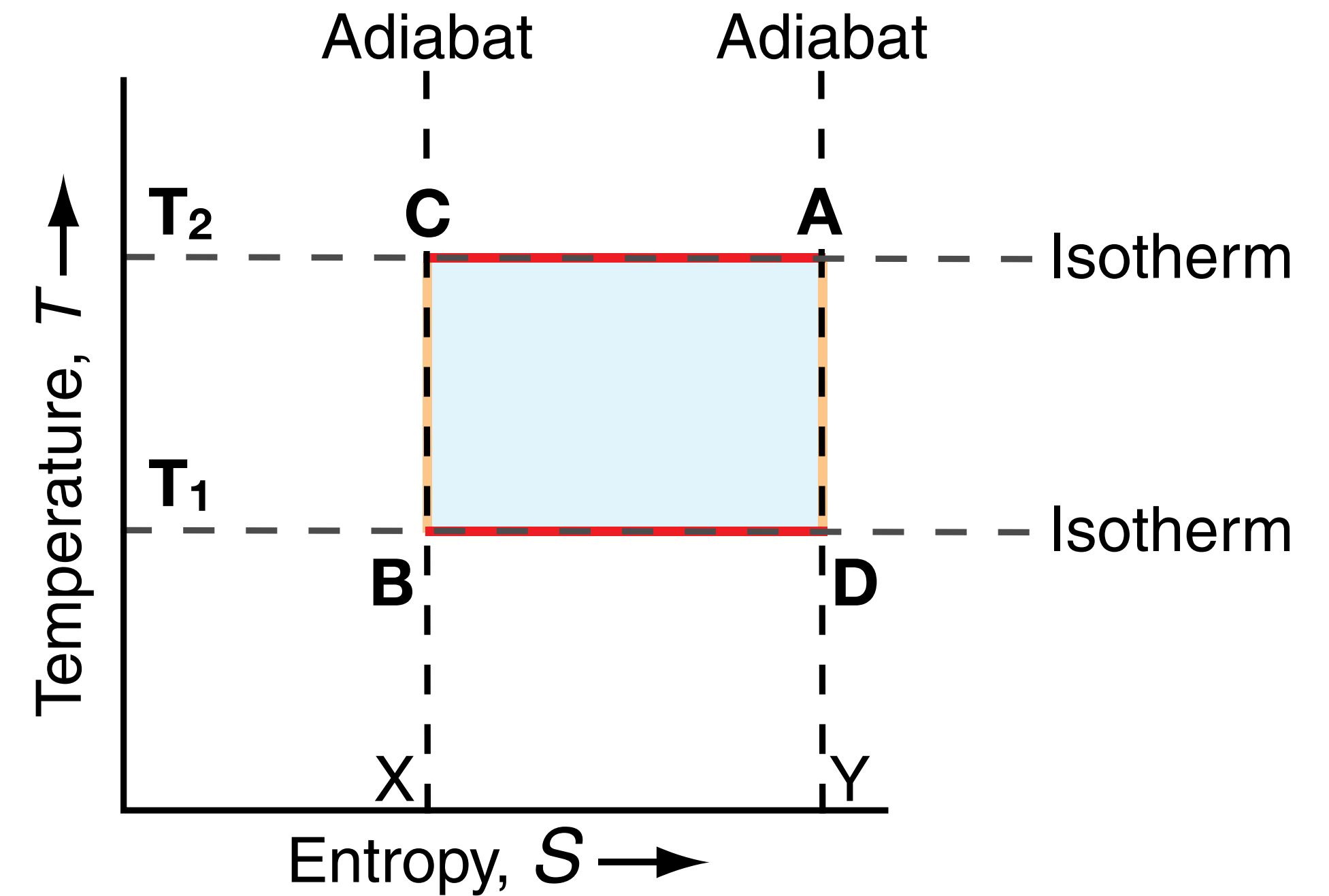
Carnot Engine

By expanding the integral into the four components of the cycle we find that

$$W = \oint p d\alpha = q_{in} - q_{out} = \varepsilon T_1 (s_{in} - s_{out})$$

$$\varepsilon = \frac{\text{Work done}}{\text{Heat absorbed}} = \frac{T_1 - T_2}{T_1}$$

Is the Carnot Efficiency



Carnot Engine

All the processes in the Carnot engine are reversible.

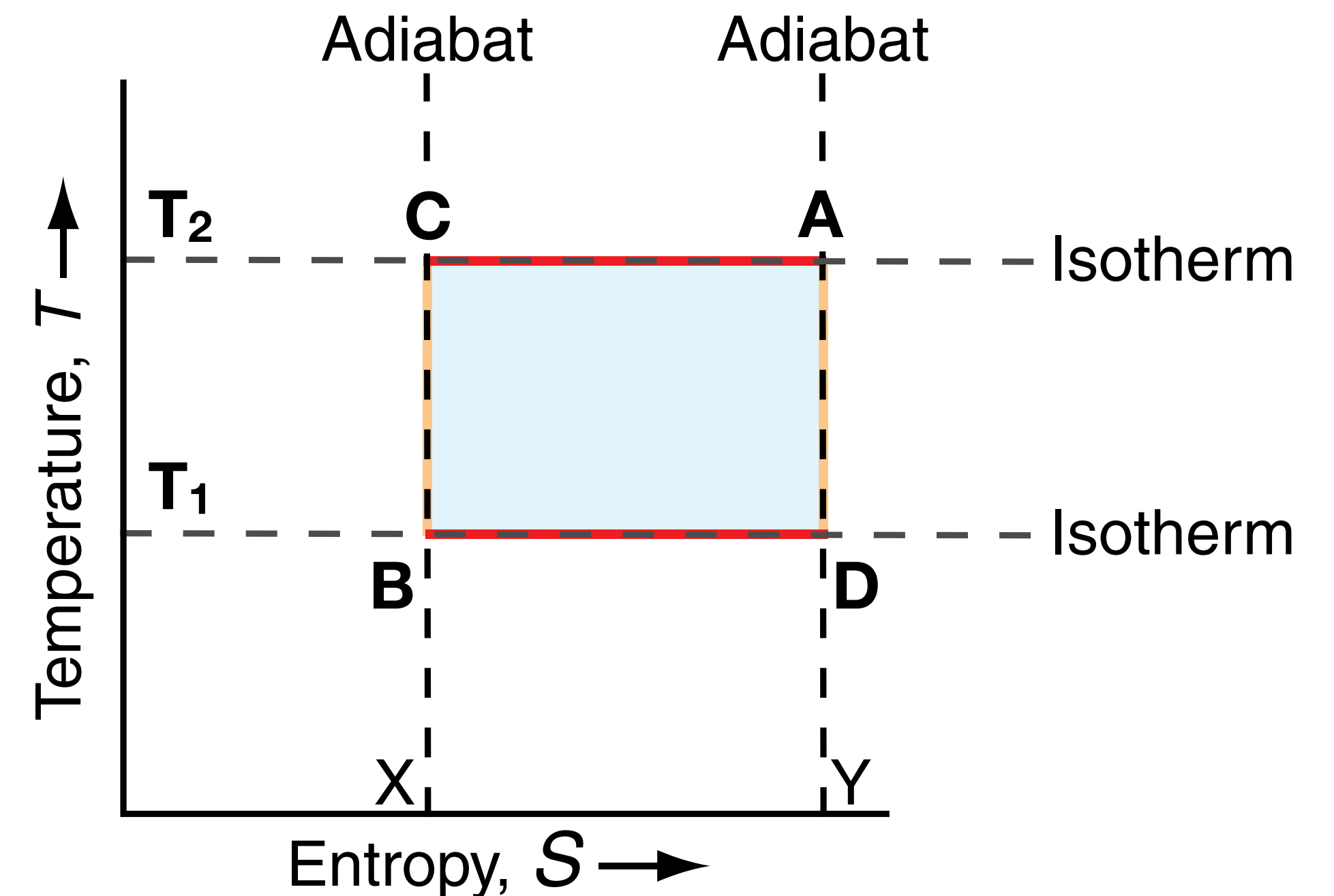
In reality, processes are irreversible, which causes additional heat loss.

Thus

$$\varepsilon = \frac{\text{Work done}}{\text{Heat absorbed}} = \frac{T_1 - T_2}{T_1}$$

Is the maximum efficiency that is possible in a given system.

$$\varepsilon(\text{Reality}) < \varepsilon(\text{Carnot})$$



We will now begin discussing the last topic of the course:

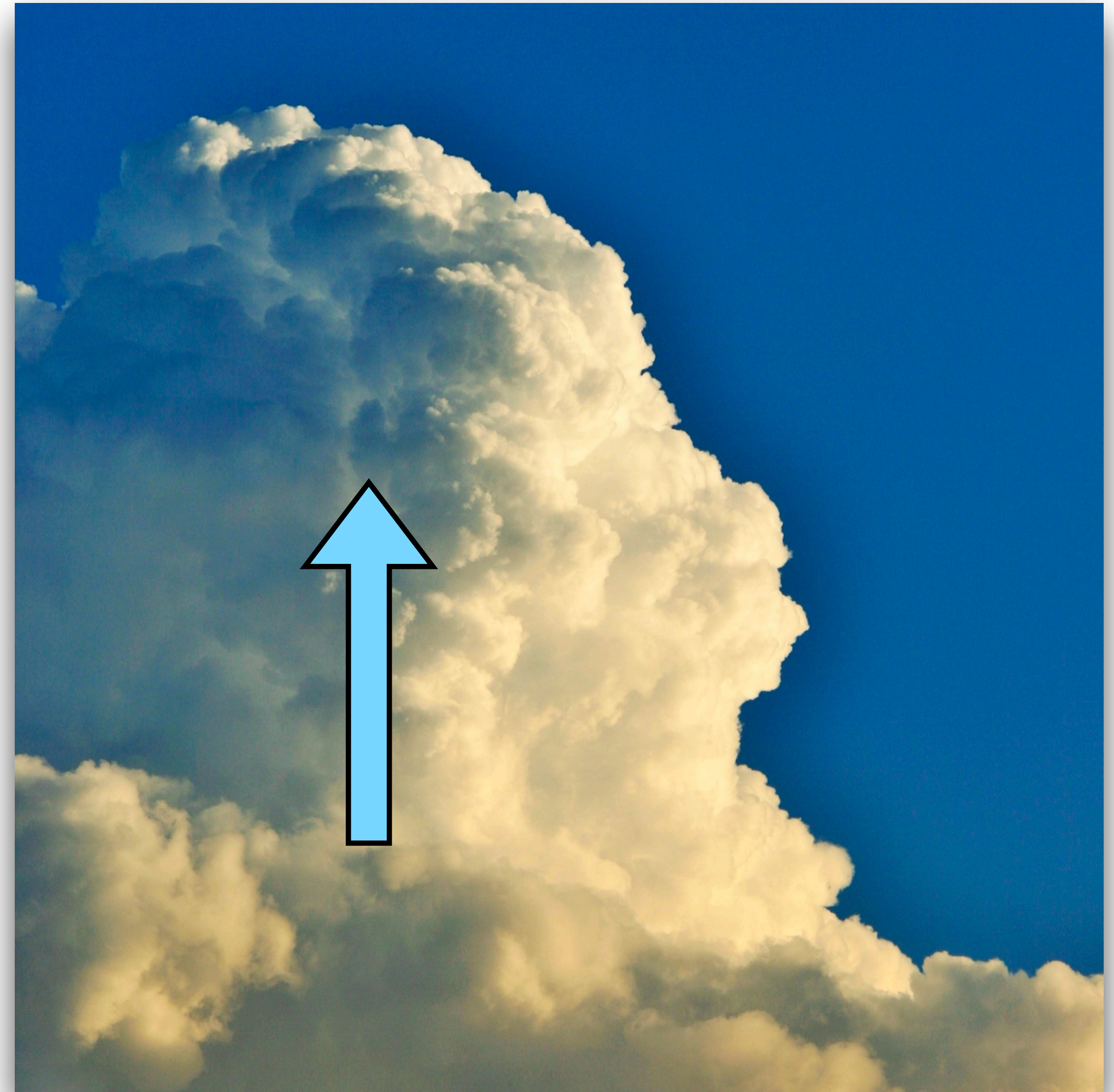
buoyancy and convection

Vertical Acceleration

Newton's second law dictates that acceleration must result from a net sum of forces.

Apply this to vertical motion

$$\frac{Dw}{Dt} = \frac{1}{m} \sum_i F_z$$



Vertical Acceleration

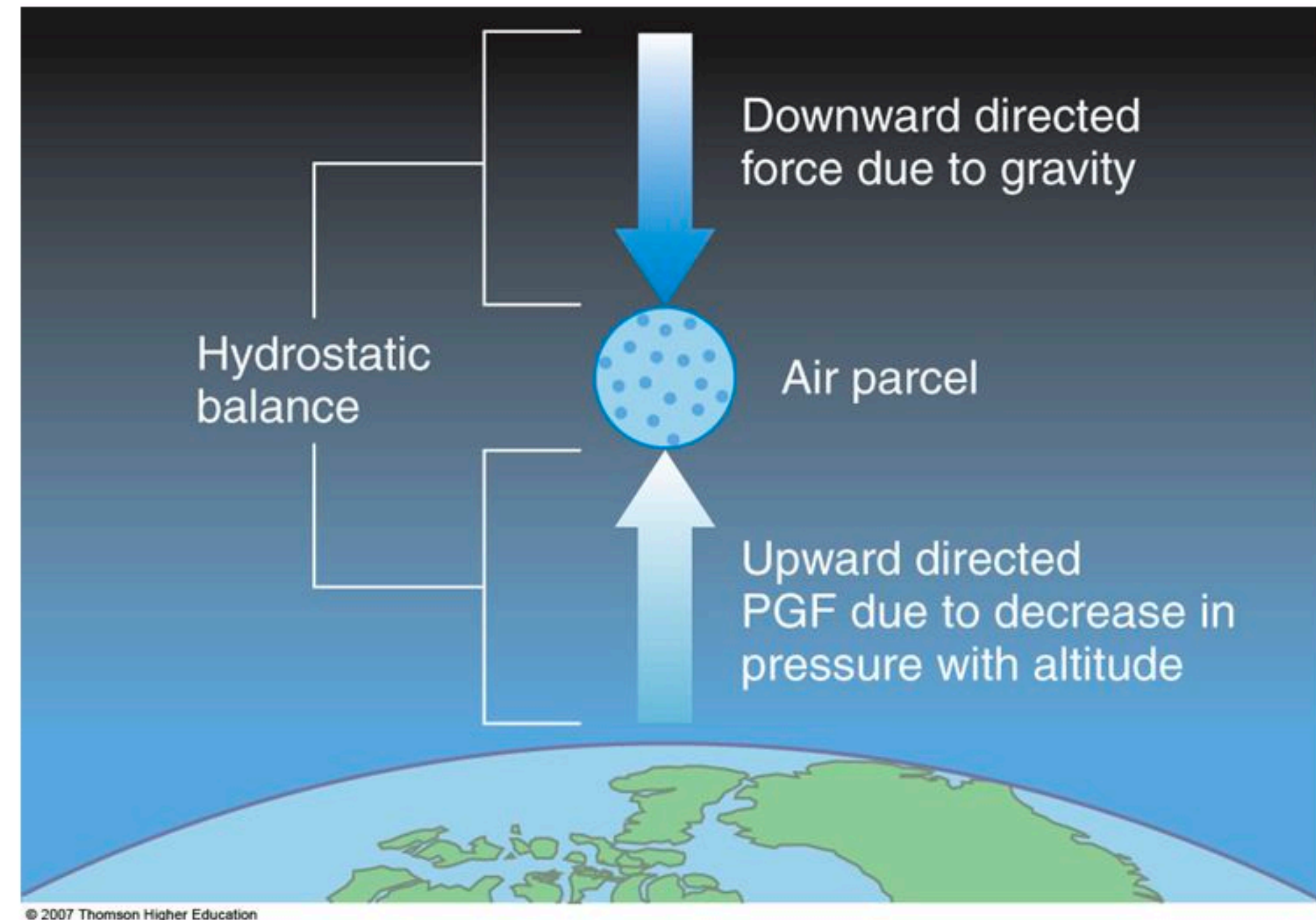
Ignoring the effects of planetary rotation and friction, the two main forces that cause vertical acceleration are gravity and the pressure gradient force.

$$\frac{Dw}{Dt} = -\alpha \frac{\partial p}{\partial z} - g$$

Acceleration

**Pressure
gradient force**

Gravity



Hydrostatic Balance

For quiescent atmospheric conditions, the atmosphere is maintained in place by a balance between the **downward gravitational force** and the **upward pressure gradient force**.

$$\rho g \simeq - \frac{\partial p}{\partial z}$$

Hydrostatic Equilibrium

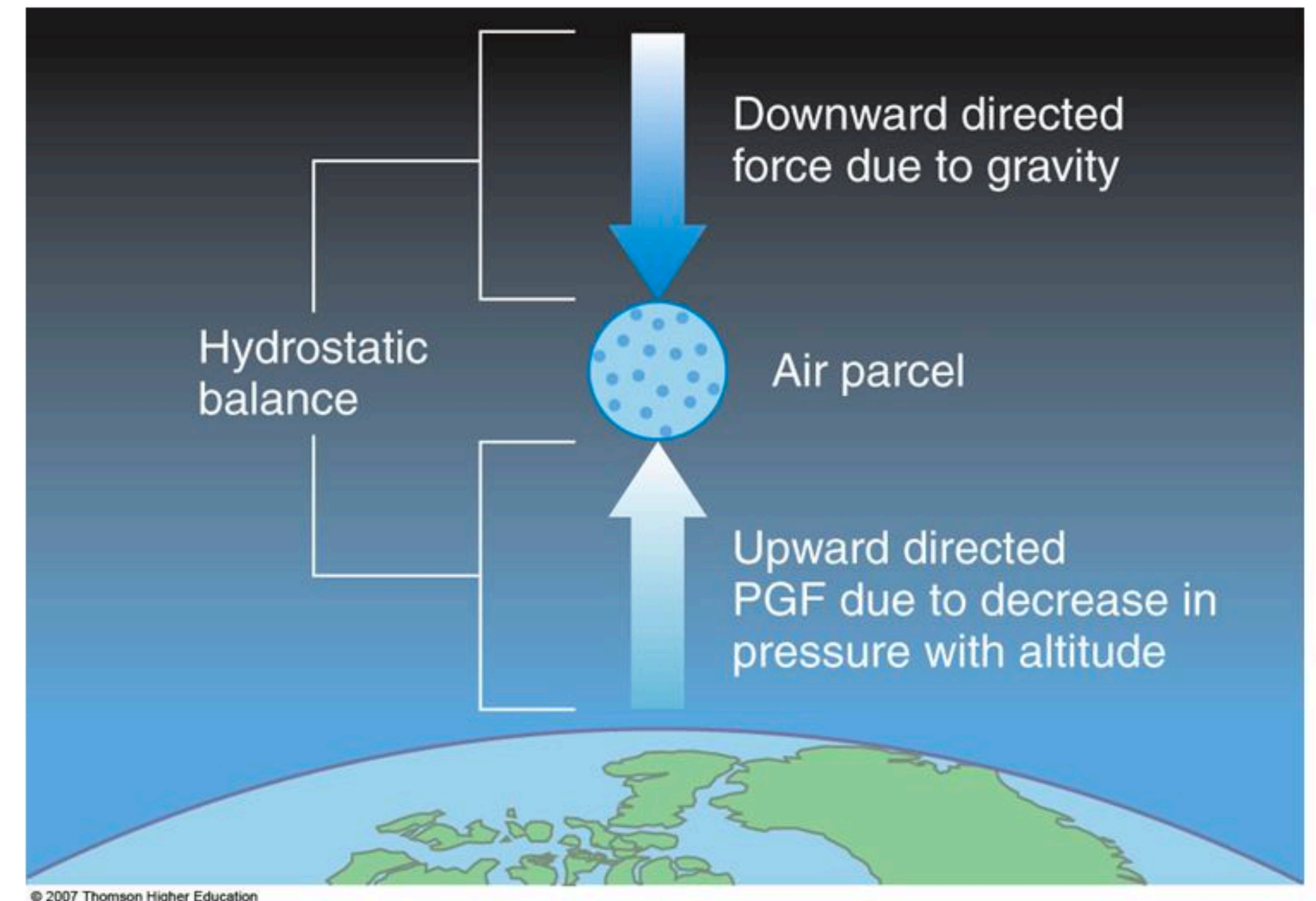


Fig. 6-13, p. 171

Defining buoyancy

We want to know how a parcel can accelerate upward. **If we assume a hydrostatic atmosphere** as the atmosphere's mean state, then

$$\rho_0 g \equiv - \frac{\partial p_0}{\partial z}$$

$$\rho = \rho_0 + \rho'$$

$$p = p_0 + p'$$

Hydrostatic Equilibrium

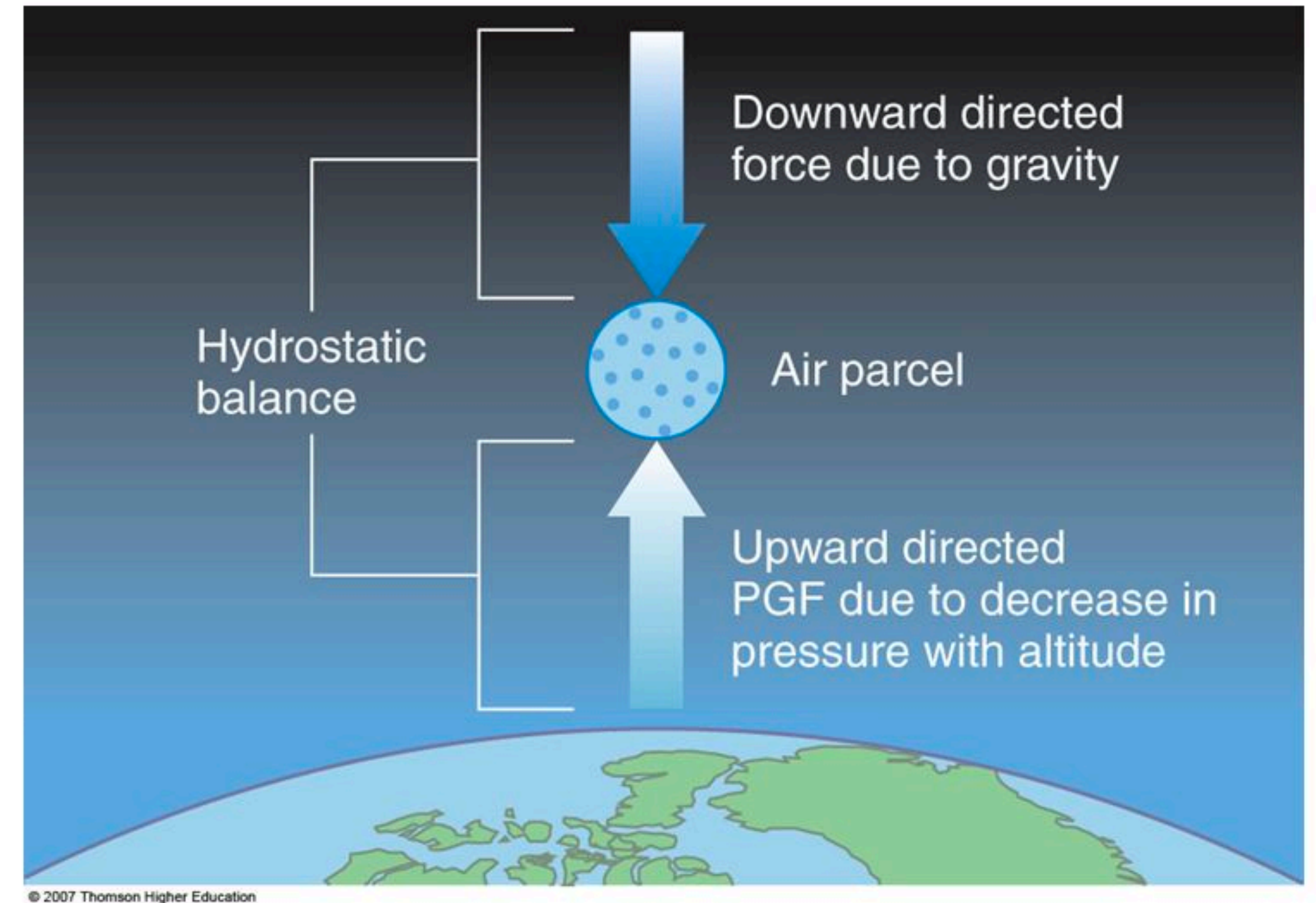


Fig. 6-13, p. 171

Defining buoyancy

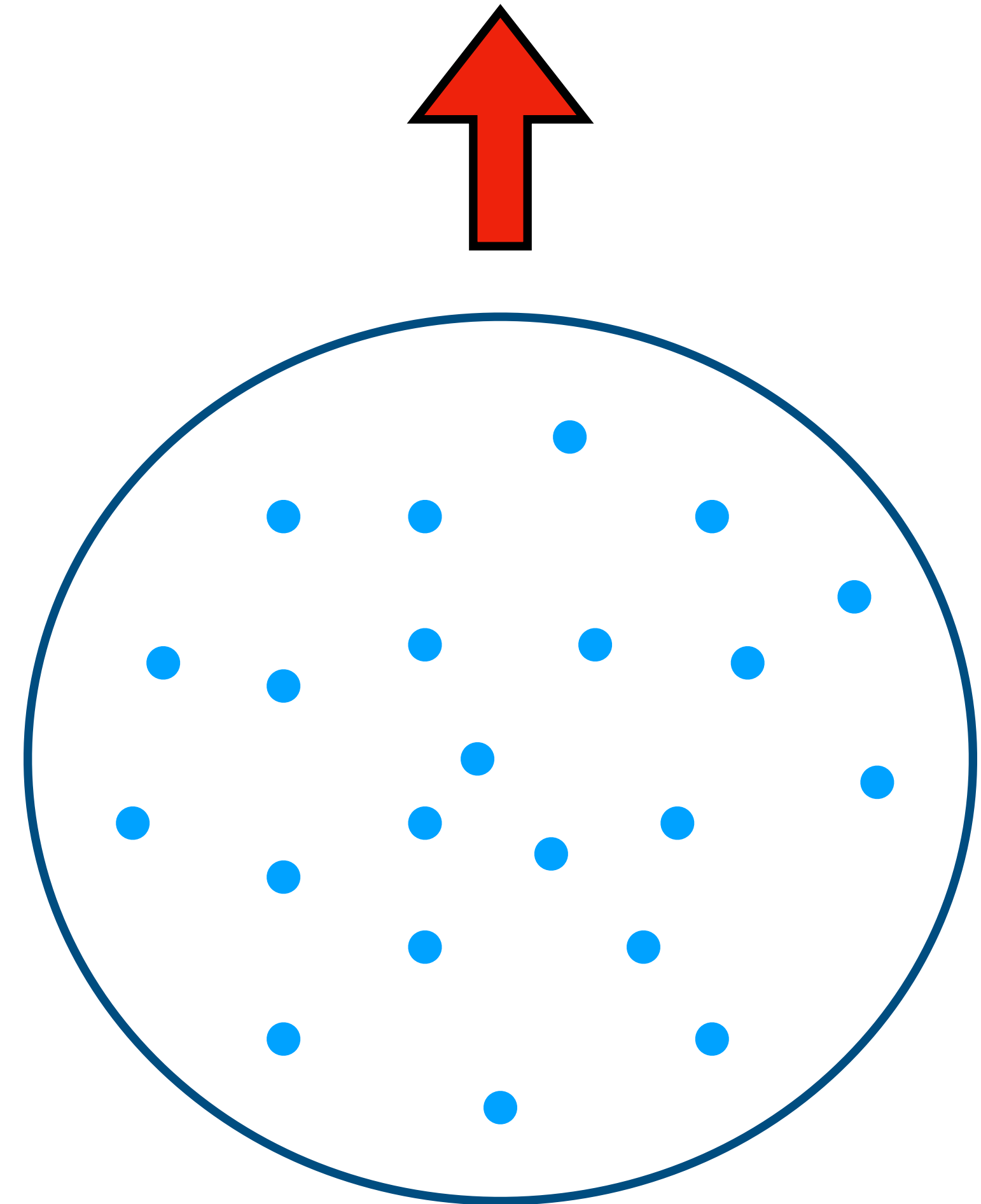
$$\frac{Dw}{Dt} = - \frac{1}{\rho_0} \frac{\partial p'}{\partial z} + B$$

Perturbation
Pressure
gradient force

Buoyancy

$$B \equiv -g \frac{\rho - \rho_0}{\rho_0}$$

How can a parcel accelerate upward?



Note: The book uses F_B for buoyancy

For moist unsaturated air

Can express the buoyancy as the difference in virtual temperature between the parcel and its surroundings.

$$B \simeq g \frac{T_v - T_{v0}}{T_{v0}}$$

