## AOS 630: Introduction to Atmospheric and Oceanic Physics Lecture 2 Fall 2021 Thermodynamic Systems and the Equation of State <br> Ángel F. Adames-Corraliza <br> angel.adamescorraliza@wisc.edu

## Last class

Concept of a "parcel": a volume of air
Large enough to make continuum assumptions (length scale large compared to the molecular mean free path).

Small enough to neglect the effects of (and feedback to) dynamics (e.g., advection).

For the purposes of thermodynamics, we consider this a "closed system" (for now).


## Atmospheric profile of density and pressure

Density and pressure decrease exponentially with height
(Note the log scale in the figure)


Wallace \& Hobbs (2006)

## Pressure in the ocean

In contrast, ocean pressure increases linearly with depth.

Parcels of water can also be defined as in air.


FIGURE 3.2 The relation between depth and pressure, using a station in the northwest Pacific at $41^{\circ} 53^{\prime} \mathrm{N}, 146^{\circ}$ $18^{\prime} \mathrm{W}$.

## Today

Define state variables

Introduce the equation of state

## Learning goals

Understand basic definitions of thermodynamics.
Answer the questions:
What is an equation of state?
Why is the equation of state important

# Supplementary reading 

Petty<br>Chapters 2 and 3

Wallace and Hobbs
Section 3.1
Ideal_Gas_Law_Derivation.pdf on Canvas

# A variable that describes the state of a system at any given time 

You do not need information about the past or future of the system

Their changes are well-defined

State variables: temperature $(T)$, density ( () , pressure (p), volume ( $V$ )

## Process Variables

## A variable that describes the transformation of a system between two states

They usually describe a path through time and/or space.

You need information about the past or future of the system

Their changes are not well-defined

Process variables: heating $(Q)$, work $(W)$

## Process vs State Variables

## State variables

Only their initial and final values matter, so that

$$
\oint_{C} d T=0
$$

And

$$
\Delta T=T_{2}-T_{1}=\int_{T_{1}}^{T_{2}} d T
$$

State variables can be written in terms of definitive integrals

## Process variables

They describe trajectories and process

$$
\oint_{C} \delta Q \neq 0
$$

And

$$
Q=\int \delta Q
$$

Process variables can only be written in terms of indefinite integrals

## Process vs State Variables

$$
\Delta T=T_{2}-T_{1}=\int_{T_{1}}^{T_{2}} d T
$$

State variables can be written in terms of definitive integrals

$$
d T
$$

Is the exact differential. It satisfies the integral above.

$$
Q=\int \delta Q
$$

Process variables can only be written in terms of indefinite integrals

Is the inexact differential. It satisfies the integral above.

## Extensive vs intensive variables

## Extensive variables

Depends on the size of the system.
Examples: Mass, Volume

## Intensive variables

Do not depend on the size of the system.

Examples: density, temperature, pressure.

These are preferred in the geosciences.

## Equation of State

An equation that describes the relationship between state variables

$$
p=f(\rho, T)
$$

Where $f$ means function

## Ideal Gases

A mixture of point particles that either don't interact with each other or have perfectly elastic collisions.

To high accuracy, our atmosphere can be described via the use of the ideal gas law
described via the use of the ideal gas law

The ideal gas law is the atmosphere's equation of state.


## Equation of state for the atmosphere

For dry air

$$
\begin{gathered}
p_{d}=\rho_{d} R_{d} T \quad R_{d}=287 \mathbf{J} \mathbf{~ k g}^{-1} \mathbf{K}^{-1} \\
\text { Dry gas constant }
\end{gathered}
$$

Partial pressure for water vapor

$$
e=\rho_{v} R_{v} T \quad R_{v}=461 \mathbf{J} \mathbf{k g}^{-1} \mathbf{K}^{-1}
$$

Water vapor constant

Partial pressure for moist air


$$
p=p_{d}+e \quad p=\left(\rho_{d} R_{d}+\rho_{v} R_{v}\right) T
$$

## Fo <br> Equation of state for seawater

$$
\begin{aligned}
& \rho=\rho_{0}\left[1-\beta_{T}\left(T-T_{0}\right)+\beta_{s}\left(S-S_{0}\right)+\beta_{p}\left(p-p_{0}\right)\right] \\
& \text { This is just an approximation. } \\
& \text { The equation that is used in } \\
& \begin{array}{l}
\text { The equation that is used in } \\
\text { ocean modeling is much more }
\end{array} \\
& \text { complicated! } \\
& \text { complicated! }
\end{aligned}
$$

$$
\begin{array}{lll}
\hline \text { Symbol } & \text { Description } & \text { Value } \\
\hline \rho_{0} & \text { reference density } & 1.027 \times 10^{3} \mathrm{~kg} \mathrm{~m}^{-3} \\
\alpha_{0} & \text { reference specific volume } & 9.738 \times 10^{-4} \mathrm{~m}^{3} \mathrm{~kg}^{-1} \\
T_{0} & \text { reference temperature } & 283 \mathrm{~K} \\
S_{0} & \text { reference salinity } & 35 \mathrm{ppt}=35 \mathrm{~g} \mathrm{~kg}^{-1} \\
c_{s 0} & \text { reference sound speed } & 1490 \mathrm{~m} \mathrm{~s}^{-1} \\
\beta_{T} & \text { thermal expansion coefficient } & 1.67 \times 10^{-4} \mathrm{~K}^{-1} \\
\beta_{S} & \text { haline contraction coefficient } & 0.78 \times 10^{-3} \mathrm{ppt}^{-1} \\
\beta_{p} & \text { compressibility coefficient }\left(=\alpha_{0} / c_{s 0}^{2}\right) & 4.39 \times 10^{-10} \mathrm{~m} \mathrm{~s}^{2} \mathrm{~kg}^{-1} \\
c_{p 0} & \text { specific heat capacity at const. pressure } & 3986 \mathrm{~J} \mathrm{~kg}^{-1} \mathrm{~K}^{-1} \\
\hline
\end{array}
$$



